



### IDC 203: INTRODUCTION TO EARTH SCIENCES





#### Physical Properties of Minerals

Property	Relation to Composition and Crystal Structure			
Hardness	Strong chemical bonds give high hardness. Covalently bonded minerals are generally harder than ionically bonded minerals.			
Cleavage	Cleavage is poor if bond strength in crystal structure is high and is good if bond strength is low. Covalent bonds generally give poor or no cleavage; ionic bonds are weak and so give excellent cleavage.			
Fracture	Type is related to distribution of bond strengths across irregular surfaces other than cleavage planes.			
Luster	Tends to be glassy for ionically bonded crystals, more variable for covalently bonded crystals.			
Color	Determined by kinds of atoms and trace impurities. Many ionically bonded crystals are colorless. Iron tends to color strongly.			
Streak	Color of fine powder is more characteristic than that of massive mineral because of uniformly small size of grains.			
Density	Depends on atomic weight of atoms and their closeness of packing in crystal structure. Iron minerals and metals have high density; covalently bonded minerals have more open packing and so have lower density.			

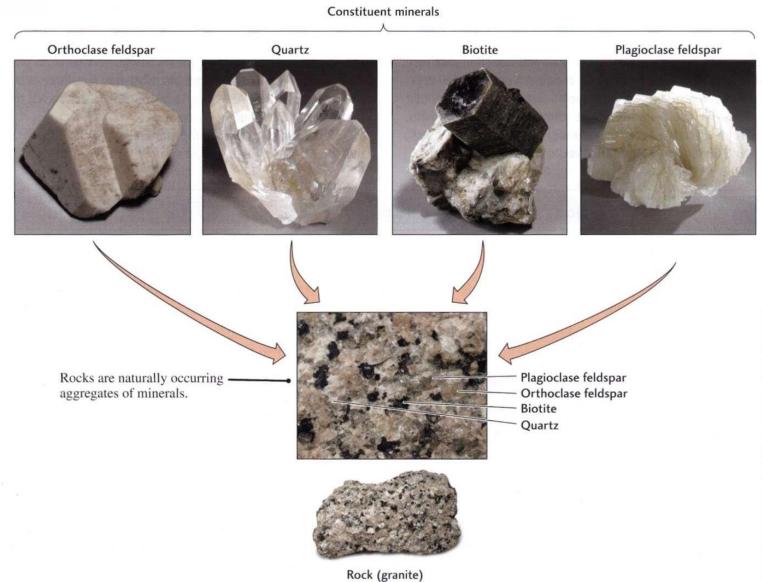
#### **Outline**

- •Rocks
- •Igneous rock- types and forms
- •Tectonic settings of Igneous rock formation
- Sedimentary rock formation

•A naturally occurring consolidated mixture of minerals or mineral-like substances

•Rocks can be also made up of non mineral matter (e.g., obsidian and coal)

•The identity of rock is determined partly by its mineral and partly by its texture.



[John Grotzinger/Ramón Rivera-Moret/Harvard Mineralogical Museum.]

#### **Texture**

•Texture describes size and shapes of a rocks mineral crystals or grains.

- The grain large enough to be seen the rock is classified as coarse grained.
- •The grain not large enough to be seen the rock is classified as fine grained.

•Rocks formed by solidification of molten rock are called igneous rock

- Weathering and erosion of rocks exposed at the surface sedimentary rocks
- •Rocks formed under high pressure and temperature in deep crust and upper mantle is metamorphic rocks

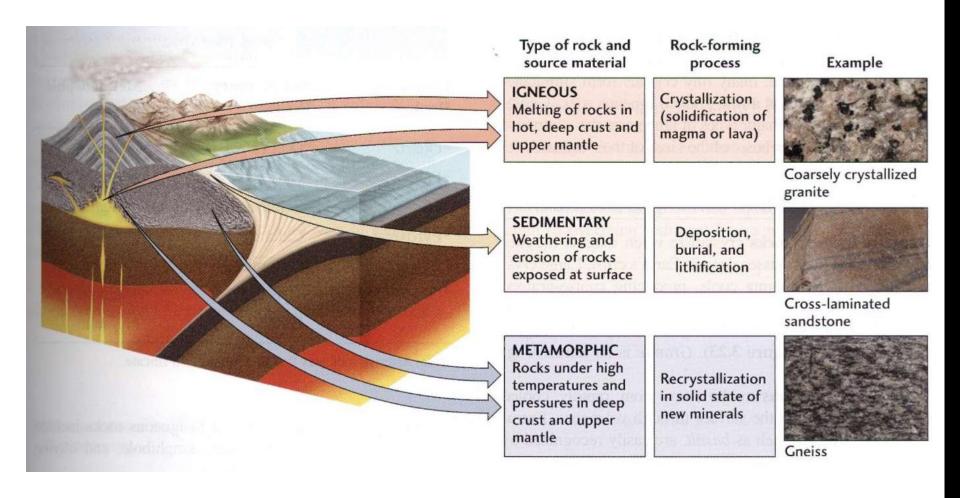


Table 3.5

Some Common Minerals of Igneous, Sedimentary, and Metamorphic Rocks

Igneous Rocks	Sedimentary Rocks	Metamorphic Rocks
*Quartz	*Quartz	*Quartz
*Feldspar	*Clay minerals	*Feldspar
*Mica	*Feldspar	*Mica
*Pyroxene	Calcite	*Garnet
*Amphibole	Dolomite	*Pyroxene
*Olivine	Gypsum	*Staurolite
	Halite	*Kyanite

Note: Asterisk indicates that the mineral is a silicate.

### **Igneous rocks**

- Formed from the cooling and consolidation of magma
- The first division of igneous rocks were made based on the evidences of texture

### Clues for early division of igneous rocks

- -Volcanic rocks
- -Laboratory studies of crystallization
- -Granite as evidence of slow cooling

### **Texture of igneous rocks**

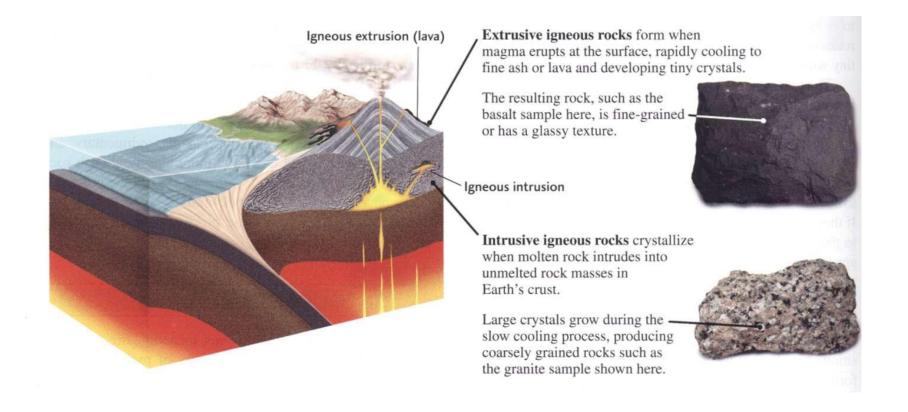
- Controlled by cooling rate
- •Grain size
- Degree of crystallinity
- Vesicularity

**Defined by texture:** 

Fine-grained: extrusive or volcanic

**Coarse-grained: intrusive or plutonic** 

#### **Igneous Rocks**



- Plutonic (intrusive) cooled below the surface
- volcanic (extrusive) cooled on the surface

# **Coarsely crystalline granite**



# Photomicrograph of granite

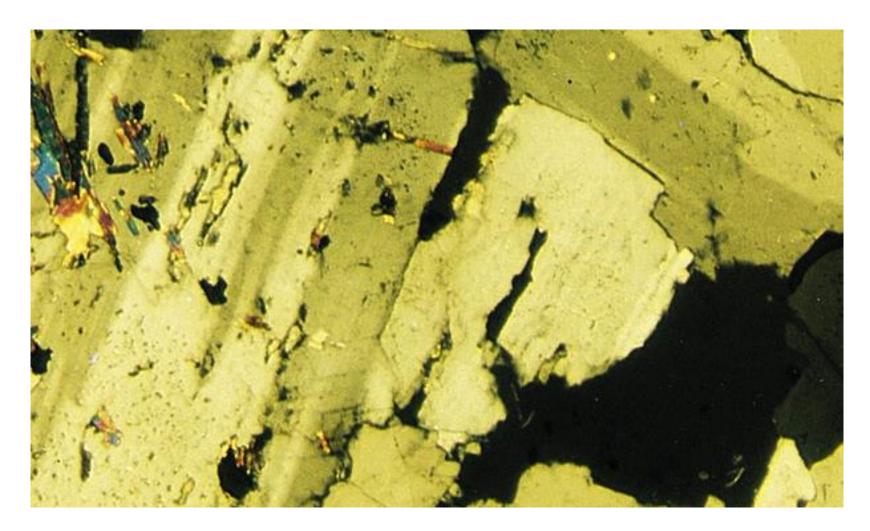


Fig. 4.1

# Finely crystalline basalt

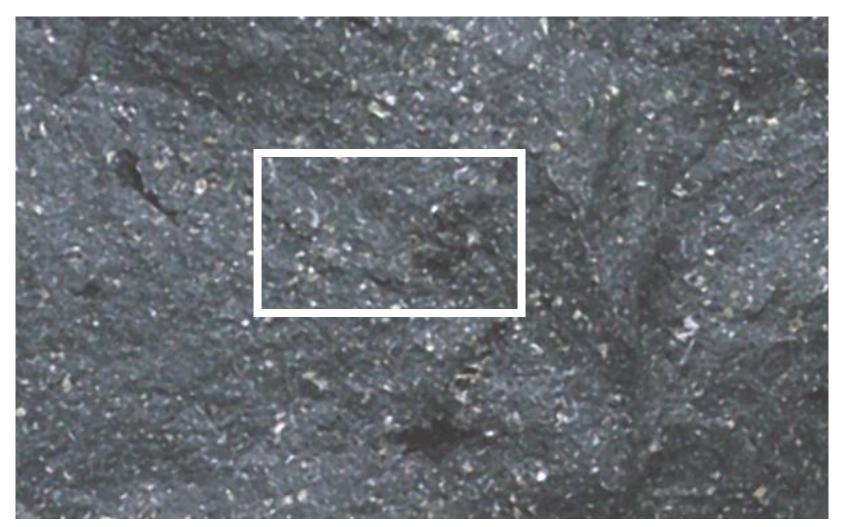


Fig. 4.1

# **Photomicrograph of basalt**

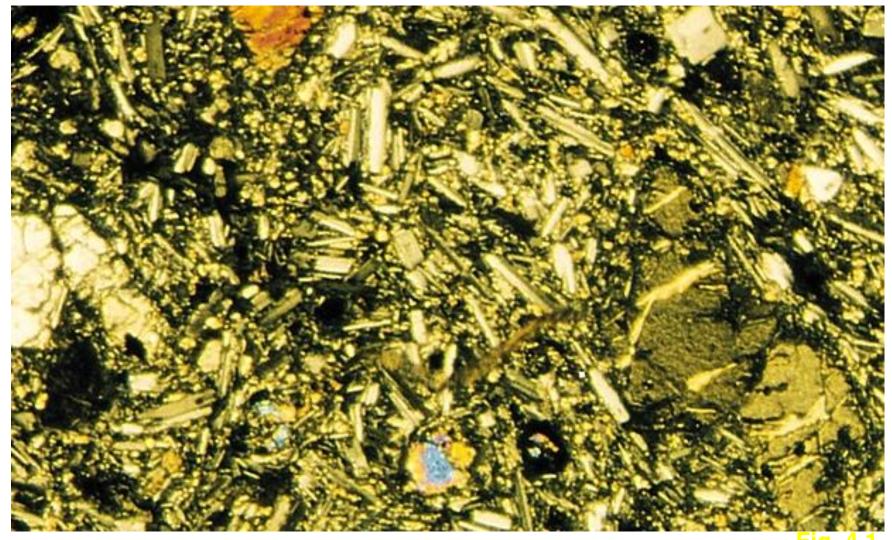


Fig. 4.1

### **Igneous Classification by texture**

Extrusive Intrusive

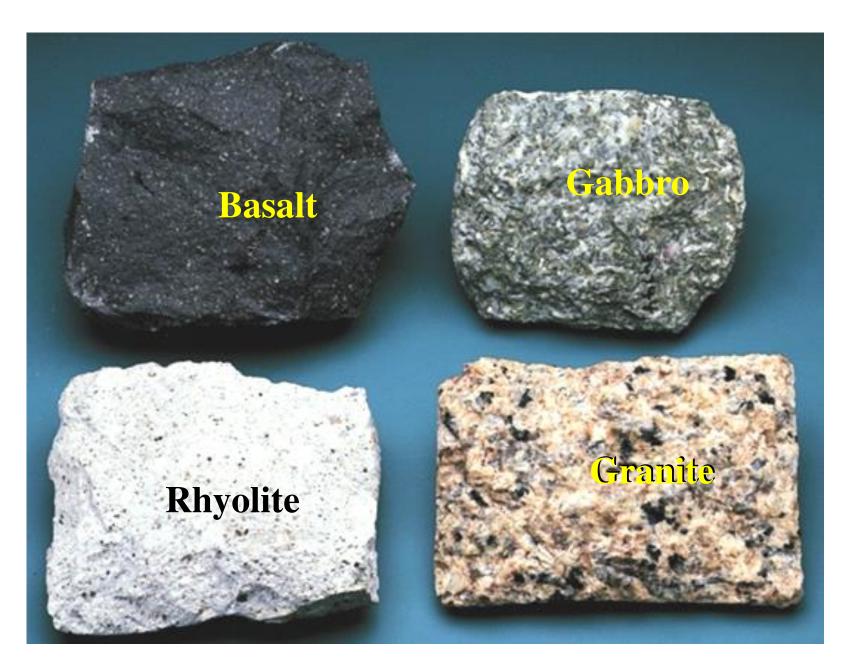
**Basalt** Gabbro

**Andesite Diorite** 

**Rhyolite** Granite

#### **Extrusive**

### **Intrusive**



## **Igneous textures**

Glassy no minerals present

Crystalline rocks made of mineral grains

Porphyritic mixture of coarse and fine

Vesicular with bubble holes

# **Igneous textures**



Glassy



**Porphyritic** 



Crystalline



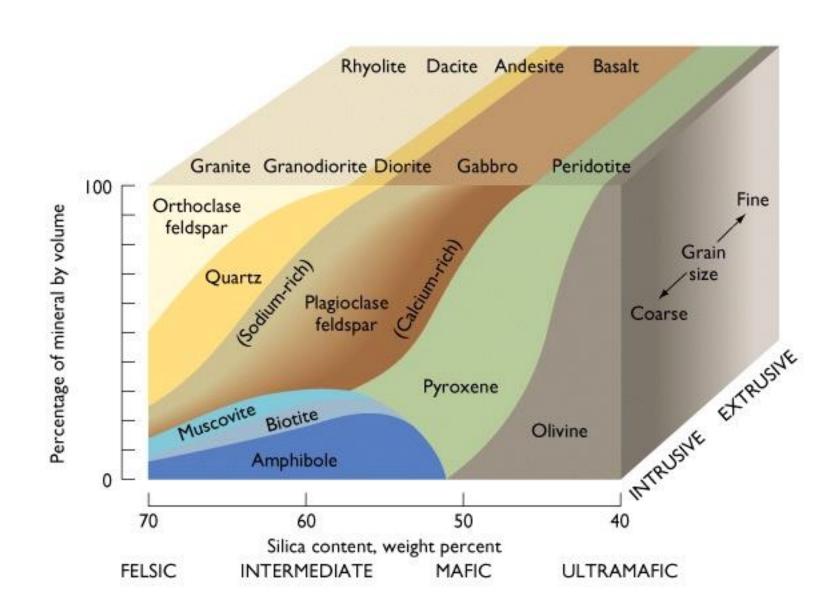
Vesicular

- Determined by composition (both chemical and mineralogical)
- Magnesium (Mg) + iron (Fe) = Mafic
- Feldspar + quartz (Si) = Felsic

Compositional Group	Mineral	Chemical Composition	Silicate Structure
	Quartz	SiO <sub>2</sub>	Frameworks
	Potassium feldspar	KAlSi <sub>3</sub> O <sub>8</sub>	
FELSIC	Plagioclase feldspar	$\left\{ \begin{array}{l} NaAlSi_3O_8 \\ CaAl_2Si_2O_8 \end{array} \right.$	
	Muscovite (mica)	$KAl_3Si_3O_{10}(OH)_2$	Sheets
	Biotite (mica)	$\left. egin{array}{c} K \\ Mg \\ Fe \\ Al \end{array} \right\} Si_3O_{10}(OH)_2$	
MAFIC	Amphibole group	$\left. egin{array}{c} Mg \\ Fe \\ Ca \\ Na \end{array}  ight\} Si_8O_{22}(OH)_2$	Double chains
	Pyroxene group	Mg Fe Ca Al SiO <sub>3</sub>	Single chains
	Olivine	(Mg,Fe) <sub>2</sub> SiO <sub>4</sub>	Isolated tetrahedra

When we talk about the chemical composition of a rock we usually speak in terms of the oxides, e.g.,

	Typical basalt	Typical granite
SiO <sub>2</sub>	50%	70%
$Al_2O_3$	15%	12%
FeO+MgO	15%	3%
CaO	8%	2%
K <sub>2</sub> O+Na <sub>2</sub> O	5%	8%



	Felsic	Intermediate		Mafic
Coarse- Grained (intrusive)	Granite	Granodiorite	Diorite	Gabbro
Fine- Grained (extrusive)	Rhyolite	Dacite	Andesite	Basalt
		Silica increasing		
		Sodium increasing		
	<b>1</b>	Potassium increasin	ıg	
		Calcium increasing		
	N	fagnesium increasi	ng	
		Iron increasing		
(Visco	osity increasi	Iron increasing	-	)

#### How do rocks melt?

Partial melting: the fracture of rock that has melted at given temperature

The temperature at the depth of crust and mantle

•Last minerals to form will melt at lowest temperature

•Biggest changes will be for small degrees of melting

#### How do rocks melt?

• The pressure due to overburden rocks also affect partial melting

• The melting temperature increases with high pressure

•Decompression melting: sudden decrease in pressure

Increase in depth in crust and mantle raises temperature

Increase in pressure raises melting range temperatures

Increase in temperature raises

Increase in temperature raises proportion of partial melt

Increase in water content lowers melting temperature

Rock composition affects melting temperature (silicic, lower temperature mafic, higher temperature)

#### **Factors affecting melting temperatures**

#### The formation of magma chamber

•The large pools of molten rock form magma chamber

•The rise of magmas at rate from 0.3 m/year to almost 50 m/year

### **Magmatic differentiation**

•The process by which rocks of varying chemical composition can arise from a uniform parent magma.

•This segregation happens in several way following a sequence described as **Bowen reaction series** 

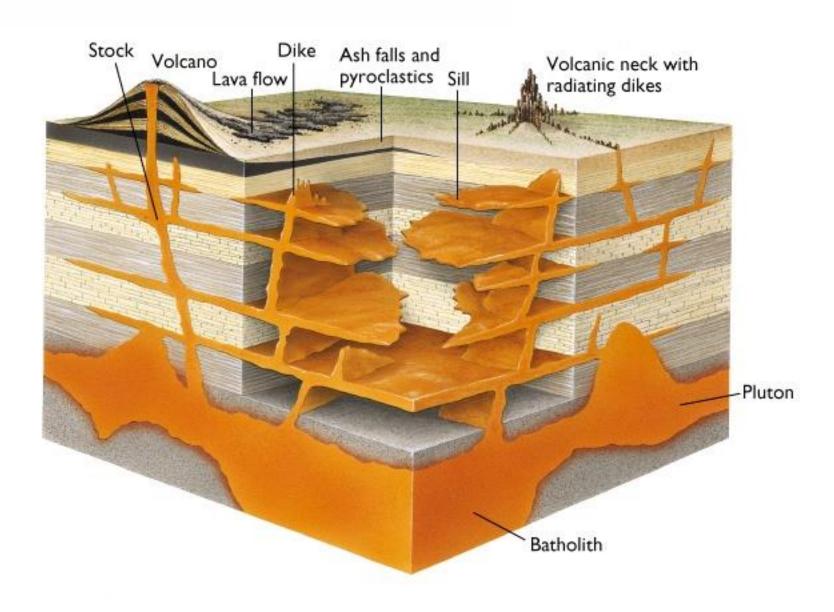
### **Fractional crystallization**

- The modification of magma by crystallization and removal of mineral phases.
- Because only certain elements will go into a given mineral, this will tend to change the composition of the remaining liquid.

#### **Bowen's reaction series**

- •Series of chemical reactions that take place in silicate magmas as they cool
- •First investigated in the 1920s and 1930s by N. L. Bowen
- •Important experiments that help us understand the evolution of magmas

# Types of igneous structures



#### Forms of intrusive igneous masses

Plutons: large igneous bodies formed in deep earth crust

Plutons can be divided into two groups:

- 1) Concordant
- 2) Discordant

Concordant: Discordant:

**Sills** Dikes

**Batholiths** 

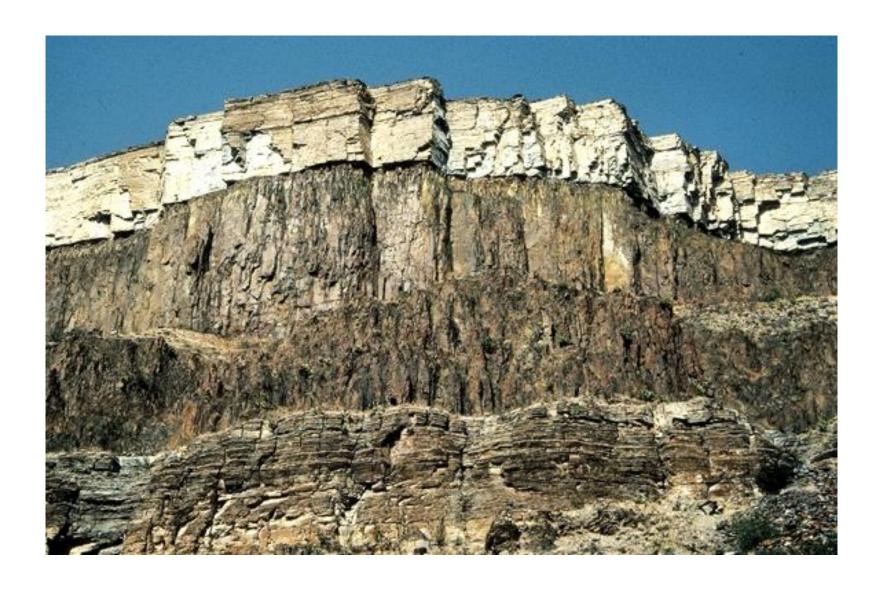
**Batholith**: Any deep-seated pluton of coarse-grained rocks that has a surface exposure of over 100 sq. km that is mostly granitic. Examples include Sierra Nevada, Coast Range, Idaho batholiths.

**Stock:** Same as a batholith, only smaller

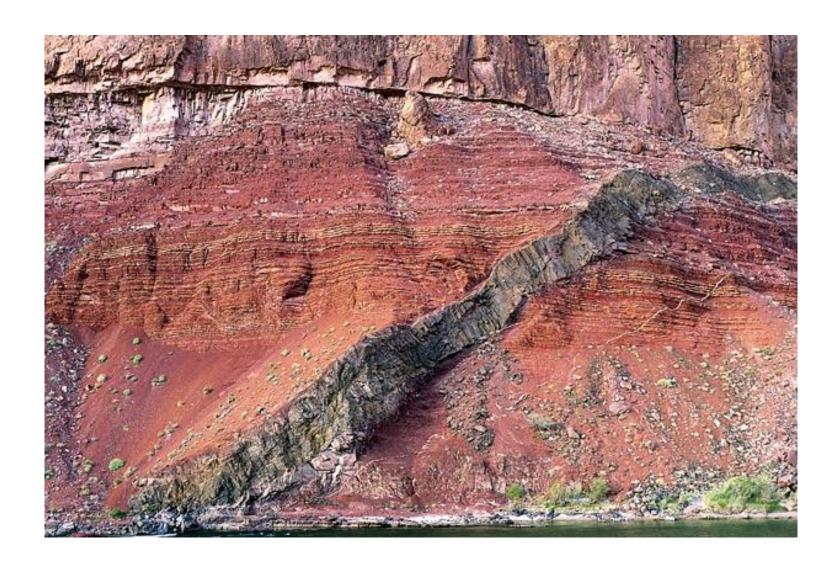
Sills: sheetlike body formed by the injection of magma between parallel layers of bedded country rock.

**Dyke**: sheetlike body formed by the injection of magma cut across the layers in bedded country rock.

## Sill



## Dike



Veins: deposits of minerals found within a rock fracture that are foreign to the country rock.

## **Veins**



Monochromatic chert cut by quartz veins

# **Pegmatite veins**



## Igneous processes and plate tectonics

#### **Sedimentary rocks**

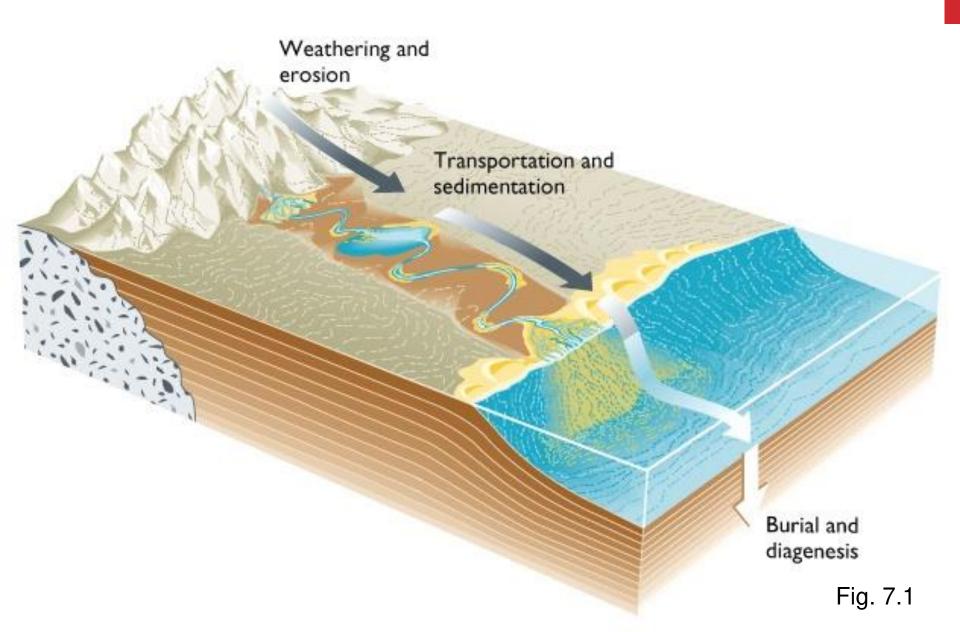
Weathering – Decomposition and disintegration of pre-existing rock into small fragments or new minerals

**Transportation** of the sediments to a sedimentary basin

**Deposition** of the sediment

**Burial and Lithification to make sedimentary rock** 

## Sedimentary stages in the rock cycle



#### **Sedimentary rocks**

**Weathering process** 

**Types of sedimentary rocks** 

**Depositional environments** 

#### Weathering process

1. Physical (Mechanical): Large rocks broken into smaller fragments with no change in composition

2. Chemical: Rocks dissolved – chemical and mineralogical composition can be altered

• new minerals may form

3. Biological: plants & animals

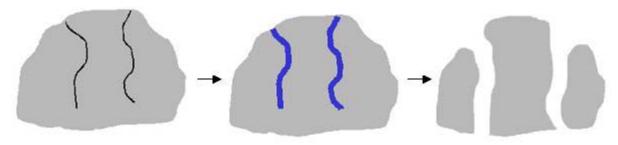
#### **Physical weathering**

1. **Joints** (also termed extensional fractures) are planes of separation on which no shear displacement has taken place



#### **Physical Weathering**

#### 2. Frost wedging



The black lines in the rock represent fractures that are occurring in the rock.

The blue lines in the rock represent water soaking into the fractures.

The water freezes and expands. If this cycle of freezing, expansion, and thawing continues, the rock will gradually disintegrate.

## **Physical Weathering**

#### 3. Salt precipitation





### **Physical Weathering**

#### 4. Abrasion

#### Physical grinding of rock fragments





**Chemical weathering** is the weakening and subsequent disintegration of rock by chemical reactions.

These reactions include

- i) Hydration
- ii) Hydrolysis
- iii) Carbonation

**Carbonation** is the process of rock minerals reacting with carbonic acid. Carbonic acid is formed when water combines with carbon dioxide. Carbonic acid dissolves or breaks down minerals in the rock.

$$CO_2 + H_2O \rightarrow H_2CO_3$$

(carbon dioxide + water → carbonic acid)

$$CaCO_3 + H_2CO_3 \rightarrow Ca^{2+} + 2HCO_{3-}$$

(calcite + carbonic acid → calcium + bicarbonate)

**Hydrolysis** is a chemical reaction caused by water.

Water changes the chemical composition and size of minerals in rock, making them less resistant to weathering.

$$2KAISi_3O_8 + 2H^+ + 9H_2O \rightarrow H_4AI_2Si_2O_9 + 4H_4SiO_4 + 2K^+$$

**Hydration** is the absorption of water into the mineral structure.

A good example of hydration is the absorption of water by anhydrite, resulting in the formation of gypsum. Hydration expands volume and also results in rock deformation.

### **Biological Weathering**

•This form of weathering is caused by activities of living organism





### Factors affect the rates of weathering process

- •Climate
- •Structural weakness
- Topography
- •Time
- Nature of rocks and minerals

#### Weathering process

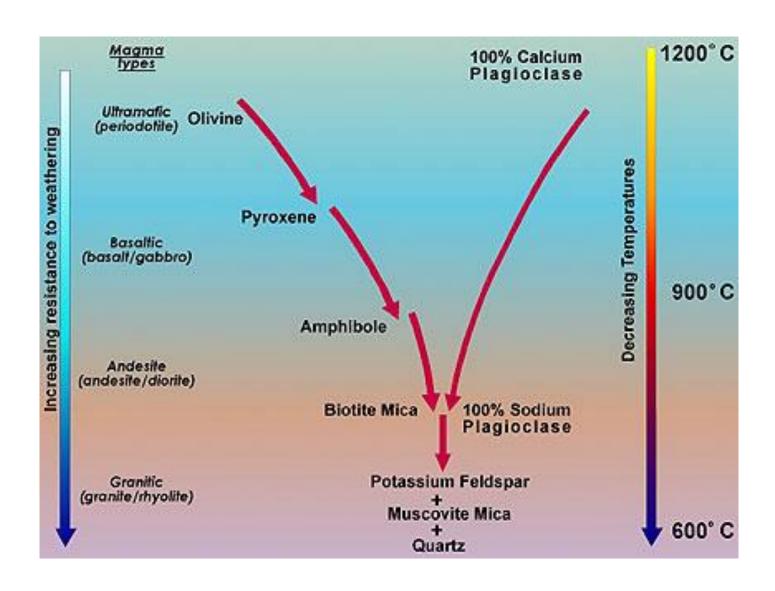


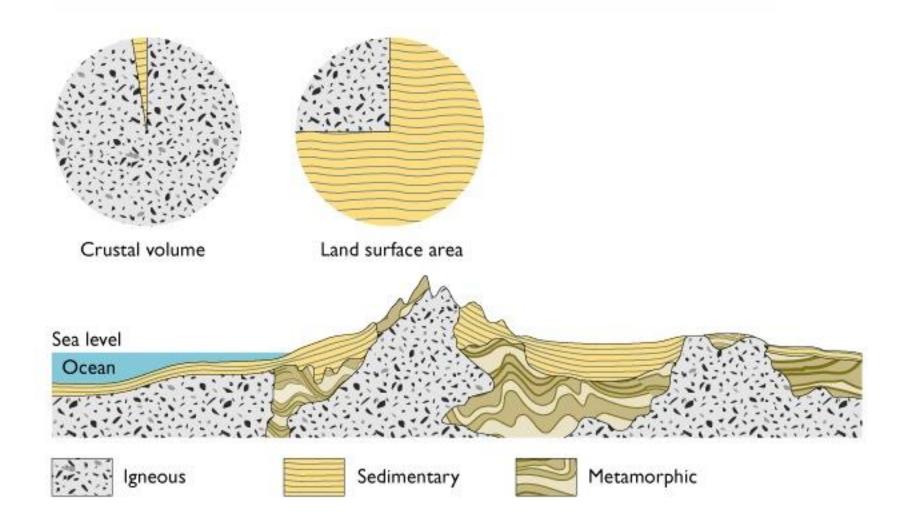
Table 7.1

## Minerals Remaining in Clastic Sediments Derived from an Average Granite Outcrop Under Varying Intensities of Weathering

## Intensity of Weathering

Low	Medium	High
Quartz	Quartz	Quartz
Feldspar	Feldspar	Clay minerals
Mica	Mica	
Pyroxene	Clay minerals	
Amphibole		

#### **HOW COMMON ARE SEDIMENTARY ROCKS?**



#### **Sedimentary Rocks**

- Detrital (clastic) rocks produced from rock fragments
- Chemical rocks produced by precipitation of dissolved ions in water
- Organic rocks produced by accumulation of biological debris,
   such as in swamps or bogs

Sedimentary rock types and *sedimentary structures* within the rocks give clues to *past environments* 

## **Transport and deposition of Clastic sediments**

•Movement of sediment by wind, ice or water.

•Mode of transport produces distinctive deposits.

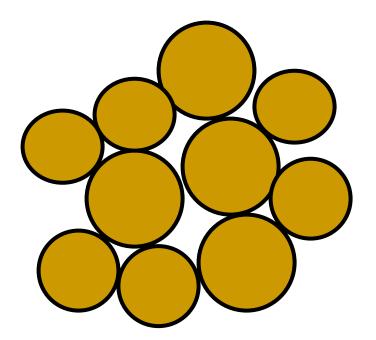
## Transport affects the sediment in several ways

**Sorting:** measure of the variation in the range of grain sizes in a clastic rock or sediment

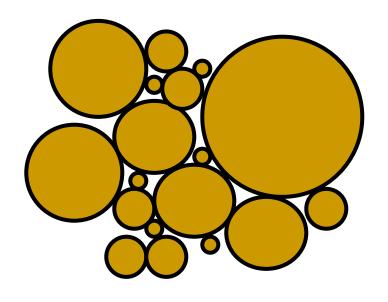
- •Well-sorted sediments indicate that they have been subjected to prolonged water or wind action.
- •Poorly-sorted sediments are either not far-removed from their source or deposited by glaciers.

## **Sorting**

**Well-sorted** 



## **Poorly-sorted**



#### Well-sorted sand



## **Poorly-sorted sand**



## Transport affects the sediment in several ways

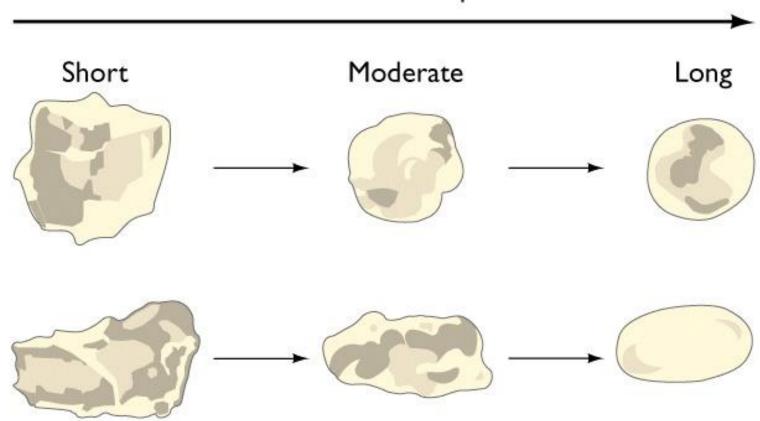
**Roundness:** measure of how rounded the corners are

Sphericity: measure of how much it is like a sphere

Sorting, roundness, and sphericity all increase with amount of transport.

# **Roundness and sphericity**

#### Distance of transport



## Types of detrital rocks

Largely based on the size of the particles, which may be anything.

- Sediment -classified by particle size
  - Boulder >256 mm
  - Cobble 64 to 256 mm
  - Pebble 2 to 64 mm ---- Gravel
  - Sand 1/16 to 2 mm
  - Silt 1/256 to 1/16 mm
  - − Clay <1/256 mm</li>

## Types of detrital rocks

Conglomerate

**Breccia** 

Sandstone (quartzite, arkose, greywacke)

**Shale** 

Mudstone

**Siltstone** 

#### **Clastic sedimentary rocks**

#### • Breccia and Conglomerate

- Coarse-grained clastic sedimentary rocks
- Sedimentary breccia composed of coarse, angular rock fragments
   cemented together
- Conglomerate composed of rounded gravel cemented together

# Conglomerate

## **Breccia**





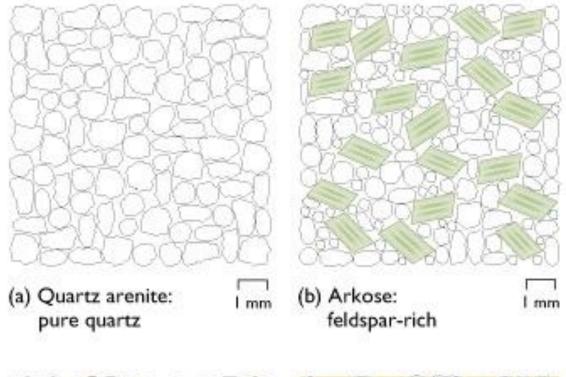
#### **Clastic sedimentary rocks**

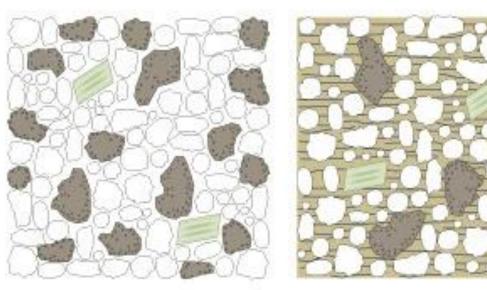
#### Sandstone

- Medium-grained clastic sedimentary rock
- Types determined by composition
  - Quartz sandstone >90% quartz grains
  - Arkose mostly feldspar and quartz grains
  - *Graywacke* sand grains surrounded by dark, fine-grained matrix, often clay-rich

# **Sandstone**







(d) Graywacke:

matrix-rich

l mm

(c) Lithic sandstone:

rock-fragment-rich

# Four major groups of sandstones

#### Shale

Shales are clastic rocks, made up mainly fine silt/clay

They are most abundant sedimentary rocks, accounts for about 80% of them

Often contain fossils

Eg. Offshore or in Lagoon

Mostly hydrous aluminum silicate in composition = from weathered feldspars

Deposition takes place under low fluvial regime or under weak water current.

• Shales are made of fine well sorted silt and clayey sediments, where normally one can expect high porosity and permeability.

## Shale



## Types of chemical sedimentary rocks

Limestone CaCO<sub>3</sub>

Chert SiO<sub>2</sub>

Salt NaCl, KCl, K<sub>2</sub>SO<sub>4</sub>

Gypsum CaSO<sub>4</sub> • 2H<sub>2</sub>O

Coal altered organic debris

#### **Chemical sedimentary rocks**

- Carbonates
  - Contain CO<sub>3</sub> as part of their chemical composition
  - Limestone is composed mainly of calcite
    - Most are *biochemical*, but can be *inorganic*
    - Often contain easily recognizable fossils

#### **Chemical sedimentary rocks**

**Limestone:** It is a non-clastic rock formed either chemically or due to precipitation of calcite (CaCO<sub>3</sub>) from organisms usually shell. These remains will result in formation of a limestone.

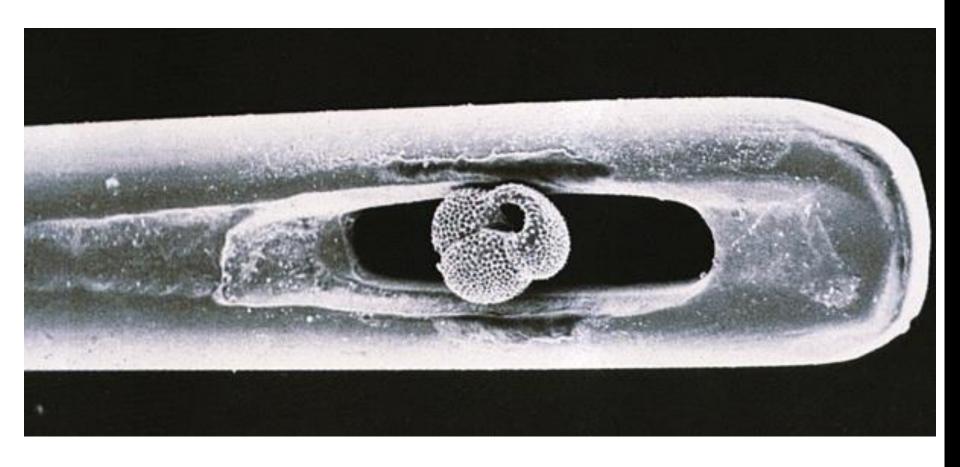
Limestones formed by chemical precipitation are usually fine grained, whereas, in case of organic limestone the grain size vary depending upon the type of organism responsible for the formation

Fossiliferous Limestone: which medium to coarse grained, as it is formed out of cementation of Shells.

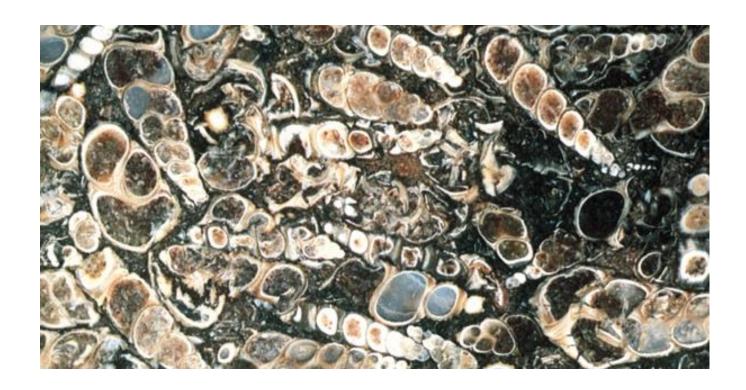
## Limestone



## Foraminifera in the eye of a needle



#### **Fossiliferous limestone**



#### **Chemical sedimentary rocks**

#### Chert

- Hard, compact, fine-grained, formed almost entirely of silica
- Can occur as layers or as lumpy nodules within other sedimentary rocks, especially limestones



## **Evaporites**

•These rocks are formed within the a depositional basin from chemical substances dissolved in the seawater or lake water

Characteristic of arid conditions

## **Evaporites**

#### Minerals precipitate according to solubility.

**Gypsum** 

50%

Halite

90%

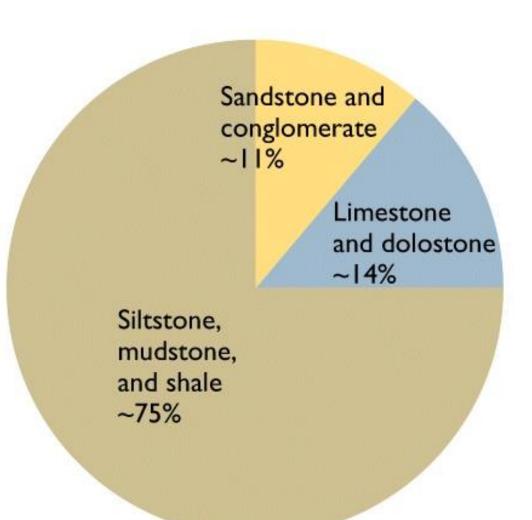


 $CaSO_4.2H_20$ 



**NaCl** 

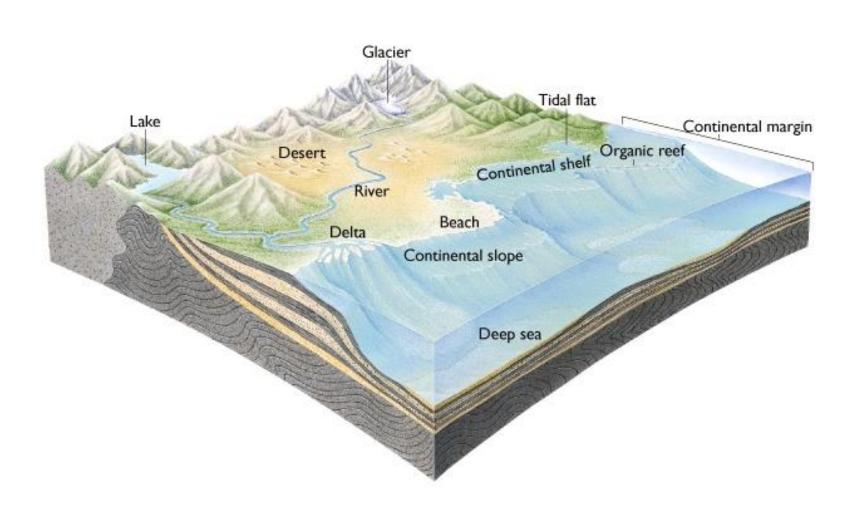
### Relative abundance of sedimentary rock types

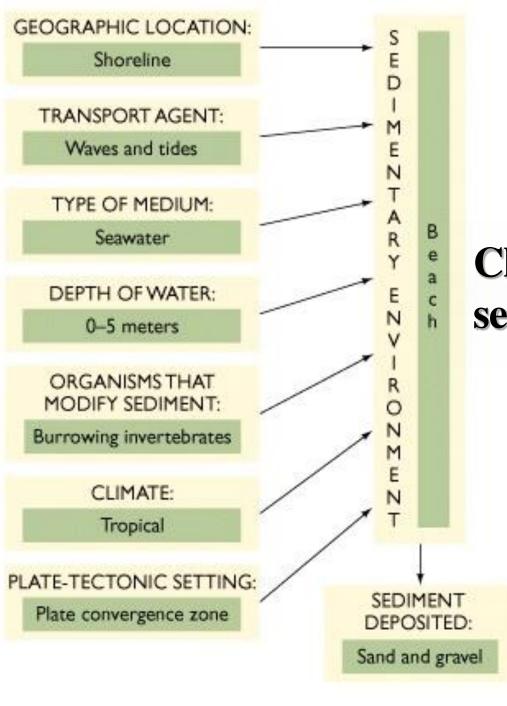


#### **Sedimentary environment**

- •Sediments accumulate in some environment of deposition or depositional environments
- •These areas receive net deposition
- •Erosion may occur, but deposition dominates
- •Features of these depositional environments are preserved in the rock record
- •Examples:
  - Sediment texture
  - •Fossils of organisms that lived in the environment
- •Ancient environments can be reconstructed from the clues that are preserved in the sedimentary rocks

# **Common sedimentary environments**





Characteristics of a sedimentary environment

## Clastic Sedimentary Environments

Environment	Agent of Transportation,	
	Deposition	Sediments
CONTINENTAL		
Alluvial	Rivers	Sand, gravel, mud
Desert	Wind	Sand, dust
Lake	Lake currents, waves	Sand, mud
Glacial	Ice	Sand, gravel, mud
SHORELINE		
Delta	River + waves, tides	Sand, mud
Beach	Waves, tides	Sand, gravel
Tidal flats	Currents	Sand, mud
MARINE		
Continental shelf	Waves, tides	Sand, mud
Continental margin	Ocean currents	Mud, sand
Deep sea	Ocean currents, settling	Mud

## Table 7.3

## Major Chemical and Biochemical Sedimentary Environments

Environment	Agent of Precipitation	Sediments
SHORELINE AND MARINE		
Carbonate (includes reef, bank, deep sea, etc.)	Shelled organisms, some algae; inorganic precipitation from seawater	Carbonate sands and muds, reefs
Evaporite	Evaporation of seawater	Gypsum, halite, other salts
Siliceous: deep sea	Shelled organisms	Silica
CONTINENTAL		
Evaporite	Evaporation of lake water	Halite, borates, nitrates, other salts
Swamp	Vegetation	Peat